

Essentials of Geology, 11e

Crustal Deformation and Mountain Building Chapter 17

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Deformation

- **Deformation** is a general term that refers to all changes in the original form and/or size of a rock body
- Most crustal deformation occurs along plate margins
- Factors that influence the strength of a rock
 - Temperature and confining pressure
 - Rock type
 - Time

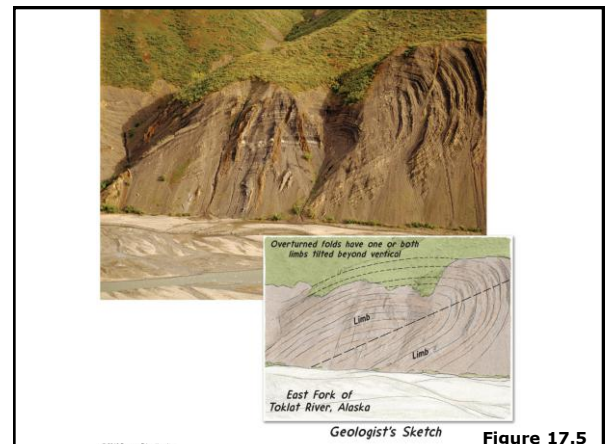
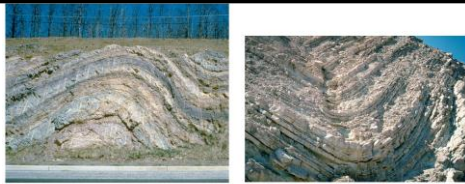
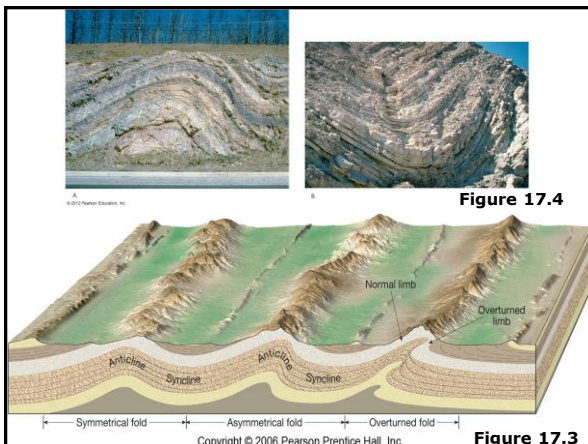
Folds

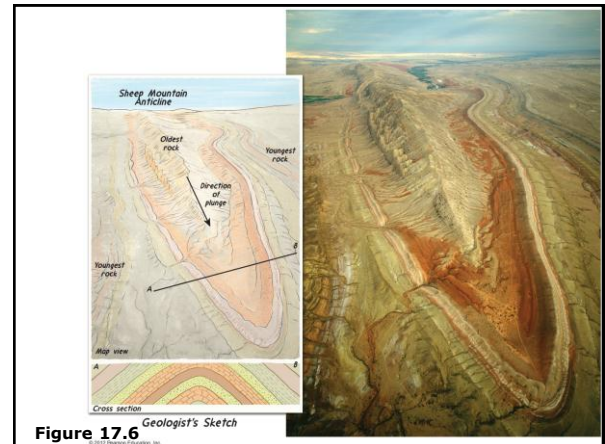
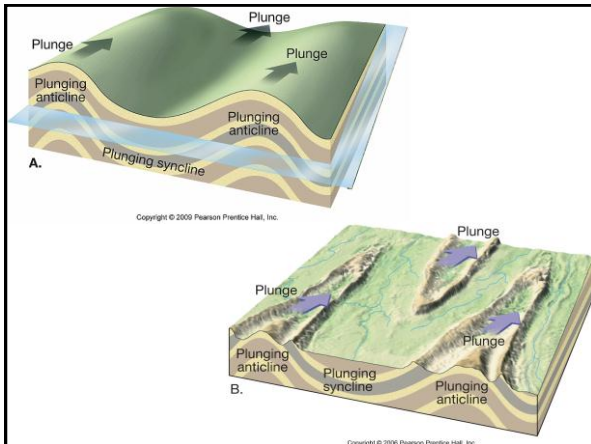
- Rocks bent into a series of waves
- Most folds result from compressional forces which shorten and thicken the crust.
- Ductile deformation.
- High temperature, soft rock, long time.
- Types of folds
 - **Anticline** – upfolded, or arched, rock layers
 - **Syncline** – downfolded rock layers

Folds

Types of folds:

- Anticlines and synclines can be
 - **Symmetrical** - limbs are mirror images
 - **Asymmetrical** - limbs are not mirror images
 - **Overturned** - one limb is tilted beyond the vertical
- Where folds die out they are said to be plunging





Folds

Types of folds:

• Other types of folds-

– Dome

- Circular, or slightly elongated
- Upwarped displacement of rocks
- Oldest rocks in core

– Basin

- Circular, or slightly elongated
- Downwarped displacement of rocks
- Youngest rocks in core

The Black Hills of South Dakota are a

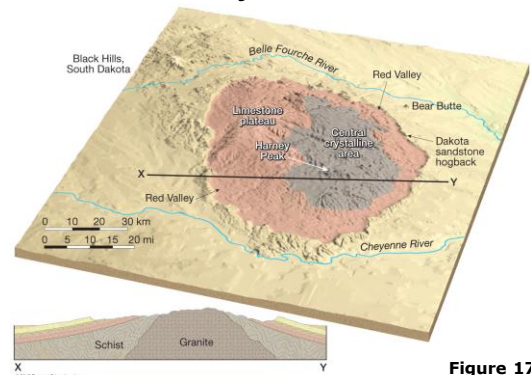


Figure 17.7

The Bedrock Geology of the Michigan Basin

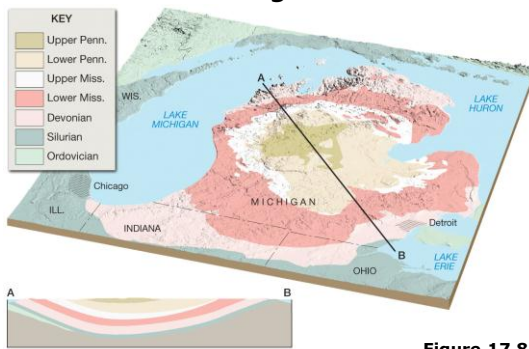


Figure 17.8

Faults

- **Faults** are fractures (breaks) in rocks along which appreciable displacement has taken place
- Brittle deformation
- Low temperature, hard rock, short time
- 2 Main types of faults (dip-slip, strike-slip)
 - **Dip-slip fault**
 - Movement along the inclination (**dip**) of fault plane
 - Parts of a dip-slip fault
 - **Hanging wall** – the rock above the fault surface
 - **Footwall** – the rock below the fault surface

Concept of Hanging Wall and Footwall Along a Fault



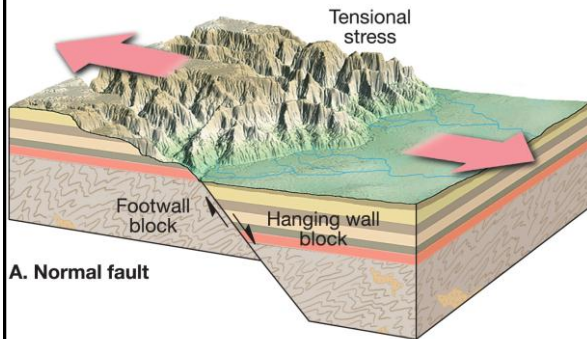
Figure 17.10

Faults

Types of faults:

- **Dip-slip fault**
 - 3 types of dip-slip faults (normal, reverse, and thrust)
 - **Normal fault**
 - Hanging wall block moves *down*
 - Associated with fault-block mountains
 - Prevalent at spreading centers
 - Caused by *tensional* forces
 - Thins and lengthens the crust

A Normal Fault



A. Normal fault

Figure 17.11 A

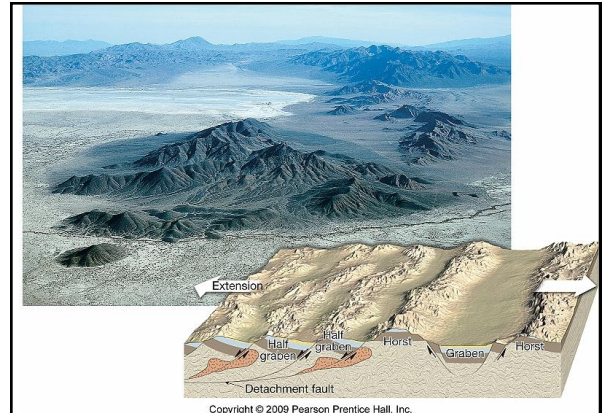
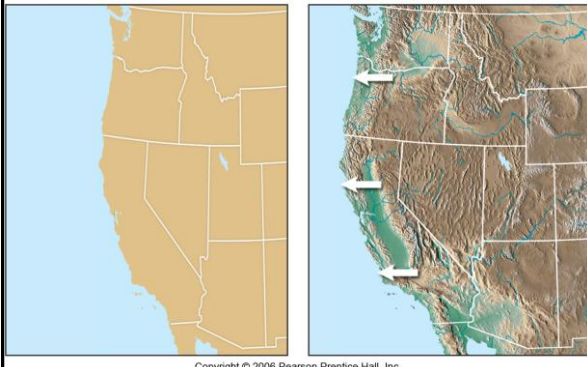


Figure 17.12

Crustal Extension

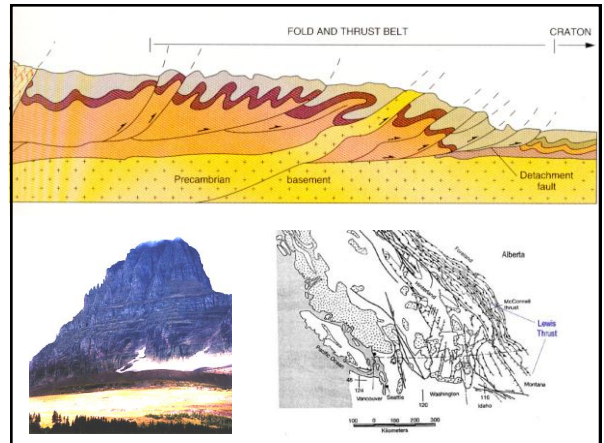
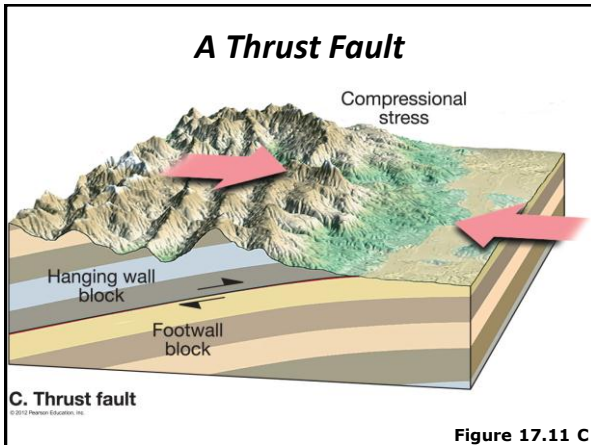
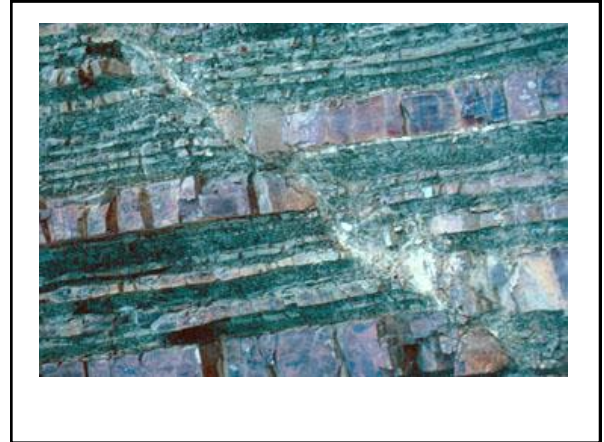
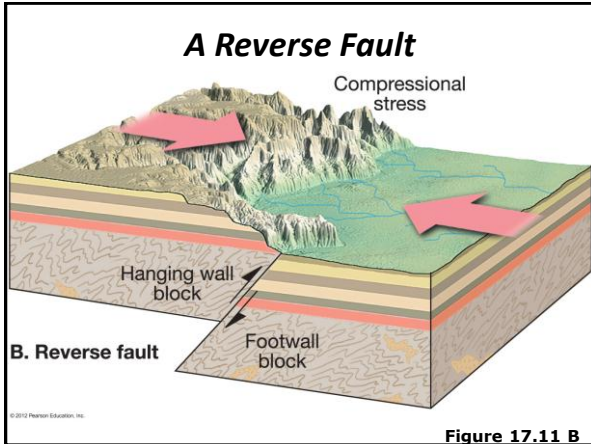


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Faults

Types of faults:

- **Dip-slip fault**
 - Types of dip-slip faults
 - **Reverse and thrust faults**
 - Hanging wall block moves *up*
 - Caused by strong *compressional* stresses
 - Thickens and shortens the crust
 - **Reverse fault** - dips greater than 45°
 - **Thrust fault** - dips less than 45°

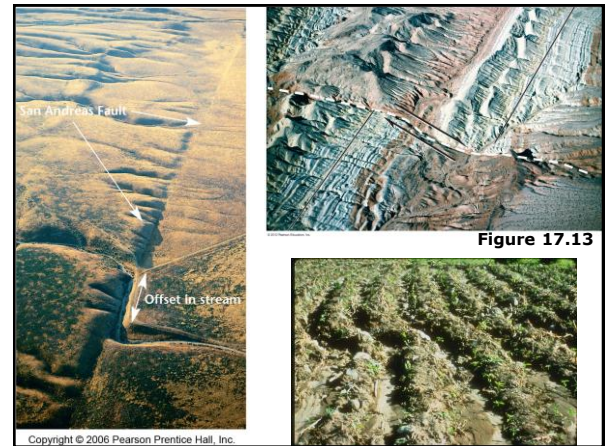
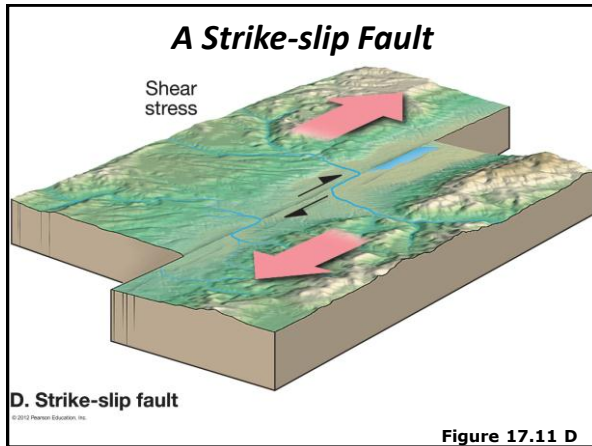


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Faults

Types of faults:

- **Strike-slip faults**
 - Dominant displacement is horizontal and parallel to the trend, or **strike**
- **Transform fault**
 - Large strike-slip fault that cuts through the lithosphere
 - Often associated with plate boundaries
 - Right or left lateral



Faults

Types of faults:

- **Joints**
 - Fractures along which no measurable displacement has occurred
 - Most are formed when rocks in the outermost crust are deformed

Figure 17.14

Mountain belts

- **Orogenesis** refers to processes that collectively produce a mountain belt.

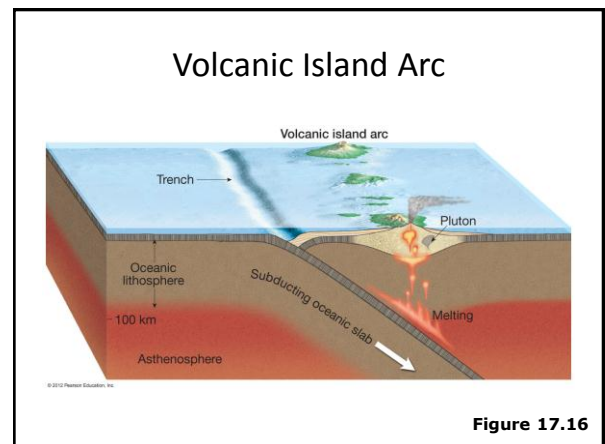
Figure 17.15

Mountain belts

Mountain building at convergent boundaries (most common):

3 subtypes... O-C C-C O-O

- **Volcanic Island Arc (O-O)**
 - Oceanic-oceanic crust convergence
 - Mafic volcanoes in the ocean
 - e.g. Aleutian Islands, AK



Mountain belts

Mountain building at convergent boundaries (most common):

- **Andean-type mountain building (O-C)**
 - Oceanic-continental crust convergence
 - e.g. Andes Mountains
 - Types related to the overriding plate
 - **Passive margins**
 - Prior to the formation of a subduction zone
 - e.g. East Coast of North America

Mountain belts

Mountain building at convergent boundaries (most common):

- **Andean-type mountain building (O-C)**
 - Types related to the overriding plate
 - **Active continental margins**
 - Subduction zone forms
 - Deformation process begins
 - **Continental volcanic arc forms**
 - **Accretionary wedge forms**
 - Examples of inactive Andean-type orogenic belts include Sierra Nevada Range and California's Coast Ranges

Orogenesis Along an Andean-type Passive Margin

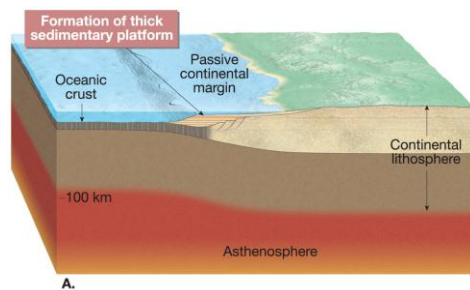


Figure 17.17 A

Orogenesis Along an Andean-type Subduction Zone

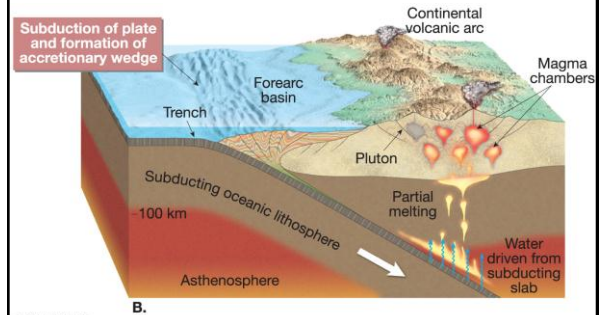


Figure 17.17 B

Orogenesis Along an Andean-type Accretionary Wedge

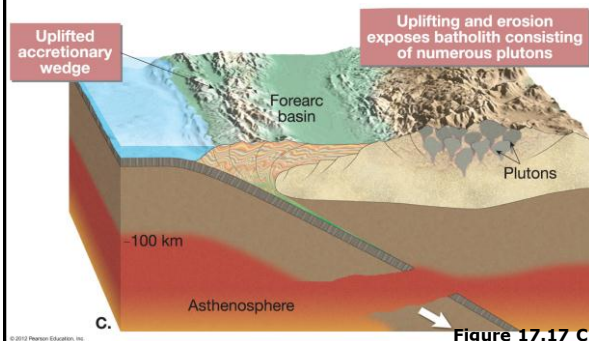


Figure 17.17 C

Mountain belts

Mountain building at convergent boundaries (most common):

- **Continental collisions (C-C)**
 - Where two plates with continental crust converge and compress.
 - e.g., India and Eurasian plate collision
 - Himalayan Mountains and the Tibetan Plateau

Formation of the Himalayas

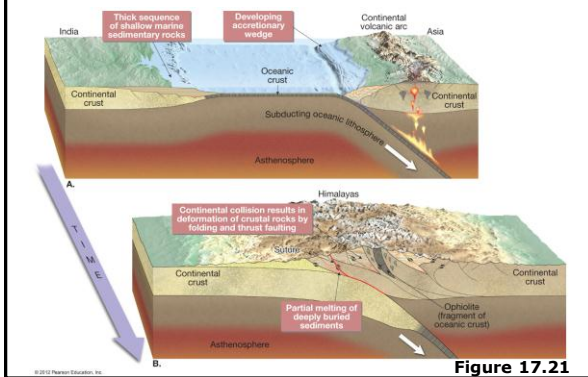


Figure 17.21

Mountain belts

Mountain building at convergent boundaries (most common):

- **Continental accretion**
 - Third mechanism of mountain building
 - Small crustal fragments collide with and accrete to continental margins
 - Accreted crustal blocks are called **terranes**
 - Occurred along the Pacific Coast

Accretion and Orogenesis

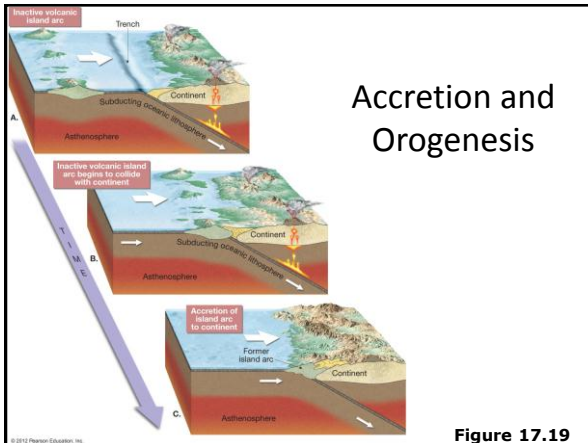


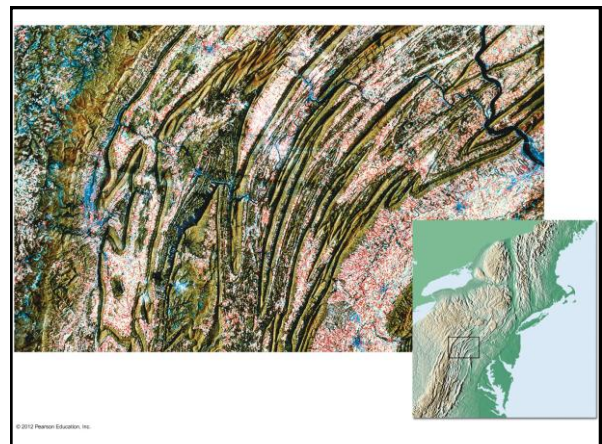
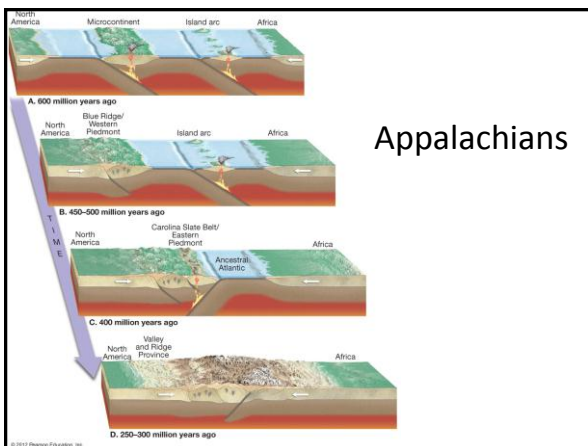
Figure 17.19

Accreted Terranes Along the Western Margin of North America



Figure 17.20

Appalachians



Mountain belts

- Buoyancy and the **principle of isostasy**
 - Evidence for crustal uplift includes wave-cut platforms high above sea level
 - Reasons for crustal uplift
 - Not so easy to determine
 - **Isostasy (2 main points)**
 - Concept of a floating crust in gravitational balance
 - When weight is removed from the crust, crustal uplifting occurs
 - Process is called **isostatic adjustment**

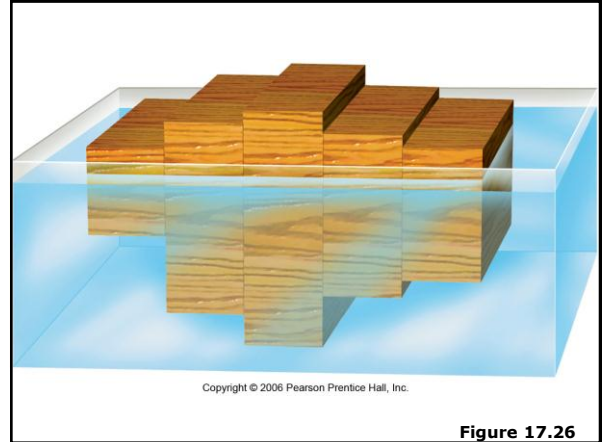
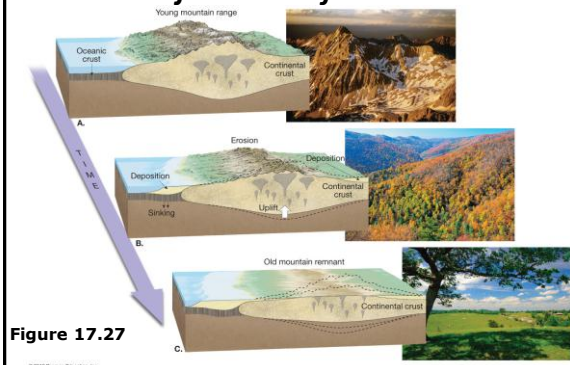


Figure 17.26

Erosion and Resulting Isostatic Adjustment of the Crust



End of Chapter 17